

# SKIDMORE

## C O L L E G E

EATON, Derek F. and NICHOLS, Kyle K., Department of Geosciences, Skidmore College, 815 N. Broadway, Saratoga Springs, NY 12866, d\_eaton@skidmore.edu

### ABSTRACT

Glacial sediments influence many surficial processes in the northeastern United States, especially landsliding and gullying. Such mass wasting is particularly common in river valleys where slopes are steep and lake clays are common. At the Miller Brook gully the glacial sediments, from the bottom up, are 1) till with a clay matrix 2) sands and gravel from near glacial runoff 3) glaciolacustrine inter-bedded clays and fine sands and 4) capping sands and gravels. Such common stratigraphies and the associated mass wasting can have significant impacts on land use in glaciated landscapes.

The exact date of gully initiation is unknown. Time series aerial photographs from the 1940s to the 1970s suggest that the gully became active in the late 1960s or early 1970s. The gully eroded due to a hydraulic conductivity difference between the sands and gravels (high) and the overlying glaciolacustrine sediments (low). Such a difference in hydraulic conductivity focused groundwater in the sands and gravels and formed a piping network that surfaced on the adjacent hillslope. In several places the pipe's roof collapsed and provided windows to the stratigraphy and the pipe. By 1998 the gully expanded to 50 m x 8 m x 2.5 m. The landowner wanted the erosion to stop and in the summer of 2001 the north bank was intentionally collapsed to fill in the gully. Since the piping network was not affected, new roof collapses were observed in the fall of 2001.

Since 2001, the gully has continued to expand; however, the gully axis is at a different orientation and is eroding naturally deposited glacial sediments, not just the fill. An August 2004 survey of the gully suggests an additional 830 m<sup>3</sup> of erosion, approximately equal to the original volume of the gully. The erosion rate since 2001 (~280 m<sup>3</sup>/y) is almost an order of magnitude higher than the estimated erosion rate from initiation until 2001 (~30 m<sup>3</sup>/y). Using the volume of an adjacent stable gully that bottoms on till is an analog, we estimate that the active gully has at least several more decades before it stabilizes. This study suggests that in order to stabilize gullies in glacial sediments, one must correctly identify the erosion processes; otherwise the resulting erosion rates could be as much as an order of magnitude greater than the natural erosion rates.

### INTRODUCTION

#### Setting

- Located on a steeply sloping terrace riser of Miller Brook near Stowe, Vermont

- Sediments deposited in glacial lake Stratigraphy (from top down)

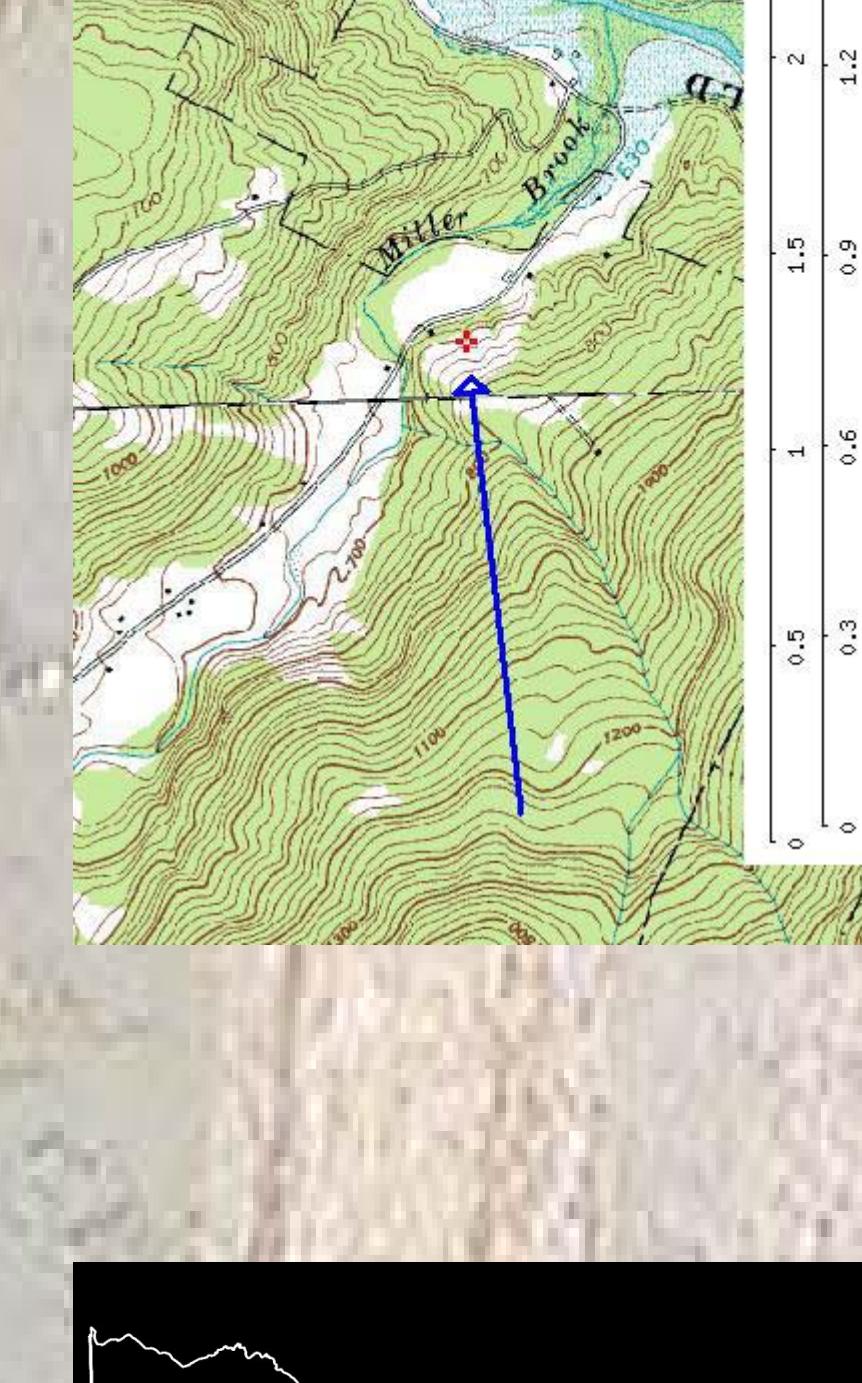
- Capping layers of sand and gravel probably associated with lowering lake levels

- Massive clays interbedded with thin layers of fine sands deposited into glacial lake

- Sand and gravel beds deposited at glacial front Adjacent gully

- Provides a natural analogue - Bottoms on till

- Stable slopes with mature trees - Alluvial fan with mature A-horizon



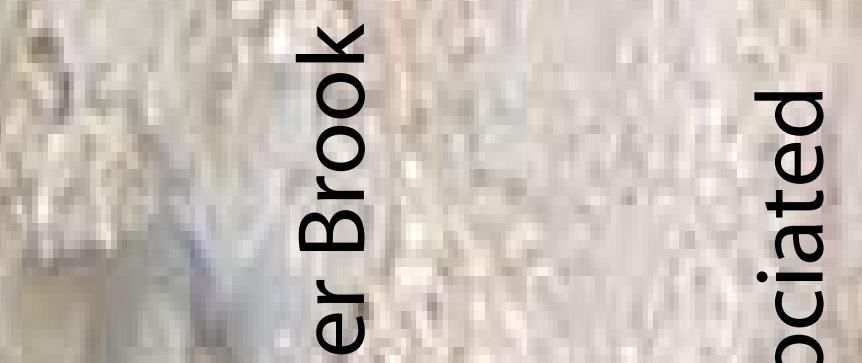
Lacustrine clays overlying till  
Scale is in 10 cm intervals



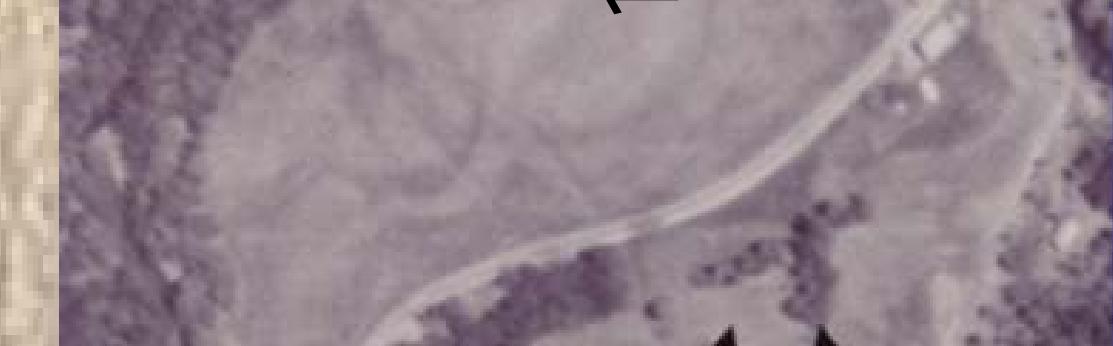
Lacustrine clays overlying till  
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Lacustrine clays overlying till  
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Lacustrine clays overlying till  
Scale is in 10 cm intervals



Top arrow in each photograph shows location of active gully  
Bottom arrow in each photograph shows location of stable gully



Top arrow in each photograph shows location of active gully  
Bottom arrow in each photograph shows location of stable gully



Top arrow in each photograph shows location of active gully  
Bottom arrow in each photograph shows location of stable gully

### Gully Processes: Piping and Mass Wasting

#### Piping

- Sands and gravels have high hydraulic conductivities ( $10^{-1}$  to  $10^{-3}$  cm sec $^{-1}$ )

- Glaciolacustrine sands and clays have low hydraulic conductivities ( $10^{-3}$  to  $10^{-7}$  cm sec $^{-1}$ )

- Sapping erodes sand (high hydraulic material) to form a pipe

- Pipe focuses water into and expands pipe network

- Eventually pipe cannot support overlying sediment and collapses forming a window

